Modelling marine interactions and dynamic ice loss in Antarctica

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Outline | Ice sheet modeling



Transantarctic Mountains at the edge of the Ross Ice Shelf. From: Bell and Seroussi (2020).

Geometry of an ice-sheet. From: Greve and Blatter (2009).

- Ice sheets \Rightarrow key role regulating the climate of the planet.
- Examples: effect on the ocean circulation through the storage and release of freshwater and high albedo.
- Conservation of mass (1), momentum (2) and energy (3), (Greve and Blatter, 2009).
- Model Yelmo (<u>https://github.com/palma-ice/yelmo</u>) ⇒
 thermomechanical ice-sheet-shelf model written in Fortran.
 Robinson et al (2020).

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0 \tag{1}$$

$$\rho \frac{d\mathbf{u}}{dt} = \mathbf{F} + \nabla \cdot \tau \tag{2}$$

$$\rho c_{\rho} \left(\frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T \right) = Q_{c} + Q_{d}$$
(3)

Summary | Research

Explore the future evolution of Antarctica and its associated uncertainty with the ice-sheet model Yelmo. Robinson et al (2020).

- Consider two periods: 2100 (comparison with ISMIP6; Seroussi et al, 2020) and
 2500 (an extension). Starting from 2015.
- Now, ensemble simulations with new forcings until **2300.**
- Other physical processes producing uncertainty.
- Implementation of a submodule for hydrology in Yelmo. Brief introduction to the model and importance related to dynamics.

IPCC, AR6 (2021)



(e) Global mean sea level change in 2300 relative to 1900

> Sea level rise greater than 15 m cannot be ruled out with high emissions

> > 9 m

8 m

7 m

Motivation | Sea level contribution



Meredith et al. (2019)

Methodology | Basal melting

Basal melt parametrization

$$B_f(x,y) = \gamma_0 \times \left(\frac{\rho_{sw} c_{pw}}{\rho_i L_f}\right)^2 \times \left(\mathsf{TF}(x,y,z_{\mathsf{draft}}) + \delta T_{\mathsf{sector}}\right) \times |\langle\mathsf{TF}\rangle_{\mathsf{draft}\in\mathsf{sector}} + \delta T_{\mathsf{sector}}|$$

Jourdain et al (2020)



Experiments

Calibration method	Forcing	$\gamma_0 ~(m/yr)$
MeanAnt Medium	ctrl	14500
MeanAnt Medium	NorESM1-M (RCP8.5)	14500
MeanAnt High	NorESM1-M (RCP8.5)	21000
MeanAnt Low	NorESM1-M (RCP8.5)	9620
PIGL	NorESM1-M (RCP8.5)	159000

MeanAnt (Mean Antarctic observations)

PIGL (Pine Island GLacier observations)

Results | Evolution until 2100



Results | Evolution until 2500 (I)



Shaded regions from Seroussi et al (2020)

Results | Evolution until 2500 (II) with PIGL

Ice Thickness Anomalies (m)

Surface Velocity Anomalies (m/yr)



Time 2015-2015



Time 2015-2015

Other physical processes | Subglacial hydrology



The 'Labyrinth' in the Dry Valleys (East Bingham (2014).



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Other physical processes | Subglacial hydrology

- From the physics, water could evolve.
- This evolution could affect ice-sheet dynamics.



How a water layer at the base impacts sliding? (Weertman, 1962)

Construction of the model

Physical principles:

- 1. Water is conserved.
- 2. Water flows from high to low values in a hydraulic potential:
 - a. Pressure
 - b. Elevation



Hewitt notes (Karthaus)

Two state variables: W (height of water) and P (pressure of the water)

Results | First test



Conclusions

- ✤ Strong dependency in the calibration of basal melting → uncertainty
- In the range of the results achieved by Seroussi et al (2020) for ISMIP6.
- In 2100, sea level rise between -12.5 mm and 19 mm. In 2500, sea level could surpass 80 cm.
- WAIS the main contributor. EAIS counterbalance effect.
- Simulations with new forcings and other GCM models until 2300.
- Subglacial hydrology is important in relation with basal sliding and dynamics.
- When fully implemented in Yelmo, we will try to assess these effects on Antarctica.

